

Evaporation Phenomenon Inside a Solar Still: From Water Surface to Humid Air

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1. Introduction

Solar stills of different designs have been proposed and investigated with a view to get greater distillate output (Murase et al., 2006). Solar stills are usually classified into two categories: a single-effect type and a multi-effect type that reuses wasted latent heat from condensation (Fath, 1998; Toyama et al., 1990). The integration between a solar collector and a still is classified into passive and active stills (Tiwari & Noor, 1996; Kumar & Tiwari; 1998). Single-effect passive stills are composed of convectional basin, diffusion, wick and membrane types (Murase et al., 2000; Korngold et al., 1996). The varieties of a still with cover cooling (Abu-Arabi et al., 2002; Abu-Hijleh et al., 1996) and a still with a multi-effect type basin (Tanaka et al., 2000) have been studied.

A basin-type solar still is the most common among conventional solar stills (Chaibi, 2000; Nafey et al., 2000; Hongfei et al., 2002; Paul, 2002; Al-Karaghoul & Alnaser, 2004; Tiwari & Tiwari, 2008). A small experimental Tubular Solar Still (TSS) was constructed to determine the factors affecting the nocturnal production of solar stills (Tleimat & Howe, 1966). Furthermore, a detailed analysis of this TSS of any dimensions for predicting its nocturnal productivity was presented (Tiwari & Kumar, 1988). They (Tleimat & Howe, 1966; Tiwari & Kumar, 1988) mainly focused on the theoretical analysis of the nocturnal production of TSS. A simple transient analysis of a tubular multiwick solar still was presented by Kumar and Anand (1992). This TSS (Tleimat & Howe, 1966; Tiwari & Kumar, 1988; Kumar & Anand, 1992) is made of heavy glass and cannot be made easily in remote areas. The cost of glass is quite high as well (Ahsan et al., 2010).

When water supply is cut off due to natural disasters (tsunamis, tornados, hurricanes, earthquakes, landslides, etc.) or unexpected accidents, a lightweight compact still, which is made of cheap and locally acquired materials, would be reasonable and practical. The second model of the TSS was, therefore, designed to meet these requirements and to improve some of the limitations of the basin-type still and of the TSS made of glass. Since

the cover material (a vinyl chloride sheet) is a little heavy and cannot form into an ideal size easily (Islam, 2006; Fukuhara & Islam, 2006; Islam et al., 2005; Islam et al., 2007a), a polythene film was adopted as a cheap new material for the cover. Consequently, the cover weight and the cost of the second model were noticeably reduced and the durability was distinctly increased. These improvements also can help to assemble and to maintenance the second model of TSS easily for sustainable use (Ahsan et al., 2010). A complete numerical analysis on TSS has been presented by Ahsan & Fukuhara, 2008; Ahsan, 2009; Ahsan & Fukuhara, 2009; Ahsan & Fukuhara, 2010a, 2010b.

Many researchers (Chaibi, 2000; Clark, 1990; Cooper, 1969; Dunkle, 1961; Hongfei et al., 2002; Malik et al., 1982; Shawaqfeh & Farid, 1995) have focused their research on conventional basin type stills rather than other types such as tubular still. Most of the heat and mass transfer models of the solar still have been described using temperature and vapor pressure on the water surface and still cover, without noting the presence of intermediate medium, i.e. humid air (Dunkle, 1961; Kumar & Anand, 1992; Tiwari & Kumar, 1988). Nagai et al. (2011) and Islam et al. (2007b), however, found that the relative humidity of the humid air is definitely not saturated in the daytime. Islam (2006) formulated the evaporation in the TSS based on the humid air temperature and on the relative humidity in addition to the water temperature and obtained an empirical equation of the evaporative mass transfer coefficient. Since the empirical equation does not have a theoretical background, it is still not known whether it can be used, when the trough size (width or length) is changed (Ahsan & Fukuhara, 2008).

In this chapter, a comparison of the evaporation and distilled water production between the first model and second one is described. Additionally, this chapter aims to present the theoretical formulation of a model for the evaporation in a TSS by dimensional analysis.

2. Production principle

The TSS consists of a transparent tubular cover and a black semicircular trough inside the tubular cover. The solar radiant heat after transmitting through a transparent tubular cover is mostly absorbed by water in the trough. Consequently, the water is heated up and evaporates. The water vapor density of the humid air increases associated with the evaporation from the water surface and then the water vapor is condensed on the inner surface of the tubular cover, releasing its latent heat of vaporization. Finally, the condensed water naturally trickles down toward the bottom of the tubular cover due to gravity and then is stored into a collector through a pipe equipped at the lower end of the tubular cover (Ahsan et al., 2010).

3. Overview of first model and second one

3.1 Structure of TSS

Fig. 1(a) shows the cross section of the second model of the TSS. The frame was assembled with six GI pipes and six GI rings arranged in longitudinal and transverse directions, respectively. The GI pipe was 0.51m in length and 6mm in diameter. The GI ring was 0.38m in length and 2mm in diameter. The reasons for selection of GI material are light weight, cheap, available in market and commonly used in different purposes. The frame was wrapped with a tubular polythene film. The film is easily sealed by using a thermal-adhesion machine (Ahsan et al., 2010).

Proper measures should be taken for disposal of such used polythene films. In Japan, a most common technique is disposed to under soil to save and keep the environment clean.

3.2 Cost of fresh water production using TSS

The most important factor for the practical application of TSS is the cost of fresh water production. The fresh water production cost using the second model is about 1245Yen/m³, which is only 13% of that of the first one. In Japan, the price of the materials is expensive. It is, therefore, expected that the water production cost will be reduced by one-third in developing and underdeveloped countries (Ahsan et al., 2010).

4. Experiment 1: method, conditions and results

4.1 Experimental method of second model

The experiment was carried out in a temperature and relative humidity controlled room to keep the external environmental conditions surrounding the TSS constant. The equipment consisted of a TSS, a solar simulator, a pyranometer (EKO, model: MS-4, $\pm 1\%$ error), a data logger (MCS, model: 486TRH, $\pm 2\%$ error), three thermo-hygrometers (VIASALA, model: HMP13, $< \pm 2\%$ error) and three electric balances (METTLER TOREDO, model: BBK422-35DLA, readability: 0.01g) connected to three computers (Ahsan et al., 2010).

The solar simulator had 12 infrared lamps (125W) arranged in six rows of two lights each. In this experiment, the temperatures of the water surface (T_w), humid air (T_{ha}), tubular cover (T_c) and ambient air (T_a), relative humidity of the humid air (RH_{ha}) and ambient air (RH_a), and radiant heat flux (R_s) were measured with thermocouples, thermo-hygrometers and a pyranometer, respectively. The measurements for T_w , T_{ha} , T_c and RH_{ha} were performed at the center of the TSS (section C-C' in Fig. 1). A thermocouple was placed in shallow water to measure T_w . Sixteen thermocouples were attached on both inner and outer surfaces of the tubular cover at eight different points at the same intervals along the circumference of the cover. The average output of these points of the inner surface was adopted as the value of T_c . A thermocouple and a thermo-hygrometer were set at 50mm below the top of the tubular cover to measure T_{ha} and RH_{ha} . The data were automatically downloaded to the data logger at one-minute intervals (Ahsan et al., 2010).

A special experimental technique to measure independently the evaporation, condensation and production of the TSS was developed. The evaporation was directly measured by placing the support frame of the trough on an electric balance, which was attached without any contact with the other components of the TSS (Fig. 1). The mass of condensation was obtained by a direct weight measurement of the TSS using a support frame on a larger electric balance. The production was directly observed by using a collector on another electric balance. The time variations of the evaporation, condensation and production were also automatically and simultaneously recorded by three computers connected to three electric balances with a minimum reading of 0.01g (Ahsan et al., 2010).

4.2 Experimental method of first model

The same experiment using the first model was carried out in the same laboratory at the University of Fukui, Japan. There was no difference in the equipment used in the first experiment and second one except an additional electric balance to observe the

condensation flux for the second one. The results of the first model were then compared with the results of the second experiment using the second model (Ahsan et al., 2010).

4.3 Experimental conditions

Table 2 summarizes the experimental conditions applied to both first and second models. The external experimental conditions were the same for both cases.

Parameter	Value
Temperature, T_a (°C)	15~35
Relative humidity, RH_a (%)	40
Radiant heat flux, R_s (W/m ²)	800
Water depth (mm)	20
Experimental duration (hr)	8

Table 2. Experimental conditions applied to both first and second models (Ahsan et al., 2010)

4.4 Experimental results

Figs. 2(a) and (b) show the time variations of the hourly evaporation flux, w_e , hourly condensation flux, w_c (for the second model only), hourly production flux, w_p , temperatures (T_w , T_{ha} and T_c) and RH_{ha} for the second model and first one, respectively. The time required for a steady state of w_e , w_c and w_p was about six hours after starting both experiments. The start of the experiment designated as $t=0$ indicates the time of switching on the solar simulator (Ahsan et al., 2010).

It can be seen from Figs. 2(a) and (b) that w_e was detected within the first hour of the experiment, while w_p was recorded two hours after the start of the experiment. There existed a big time lag between w_e and w_p . However, the time lag between w_e and w_c was very small and it was hard to distinguish the difference between them in Fig. 2(a) (Ahsan et al., 2010).

It was found that w_e and w_p gradually decreased in both models as T_a fell from 35 to 15°C. The values of w_e and w_p were slightly lower in the second model than in the first one under the same experimental conditions. The drop in the values of w_e and w_p would be a result of the difference in the design of the first model and second one. It was observed that there was an obstruction of the trickle down of the condensed water on the polythene film due to the GI pipes, horizontally arranged inside the cover of the second model as shown in Fig. 1(a). This obstruction might be the cause of less condensation and production rate for the second model of TSS (Ahsan et al., 2010).

A further important point seen in Fig. 2 is that RH_{ha} was remarkably below 100% in both models, i.e. the humid air was definitely not saturated. If the vapor density of the humid air, ρ_{vha} is saturated, the evaporation condition on the water surface, i.e. $\rho_{vw} > \rho_{vha}$ (ρ_{vw} : vapor density on the water surface) is not satisfied, because of $T_{ha} \geq T_w$ (see Fig. 2(a)) (Ahsan et al., 2010). Nagai et al. (2002) reported the same result from their experiment using a basin-type still.

Since the humid air is definitely not saturated, it is inferred that w_e , w_c and w_p would be strongly affected by the humid air temperature and relative humidity fraction, T_{ha}/RH_{ha} . Fig. 3 shows the relationship of w_e , w_c and w_p with T_{ha}/RH_{ha} for the first model and second one. It is found that $w_p \approx w_c \approx w_e$ and these (w_e , w_c and w_p) were proportional to T_{ha}/RH_{ha} , regardless of the models (Ahsan et al., 2010).

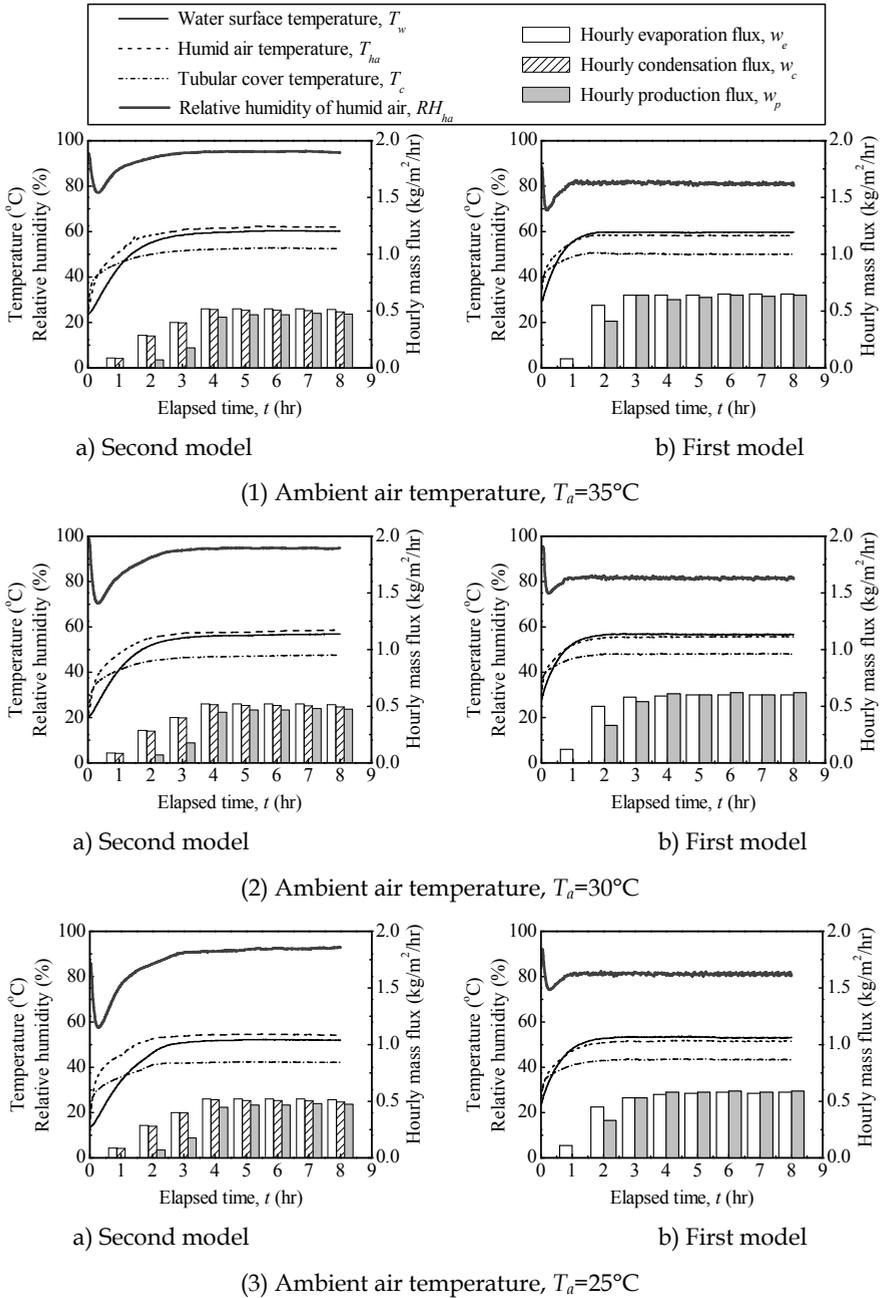


Fig. 2. Time variations of the hourly evaporation flux, w_e , hourly production flux, w_p , temperatures (T_w , T_{ha} and T_c) and RH_{ha} for different T_a ranged from 15 to 35°C for the first model and second one (Ahsan et al., 2010)

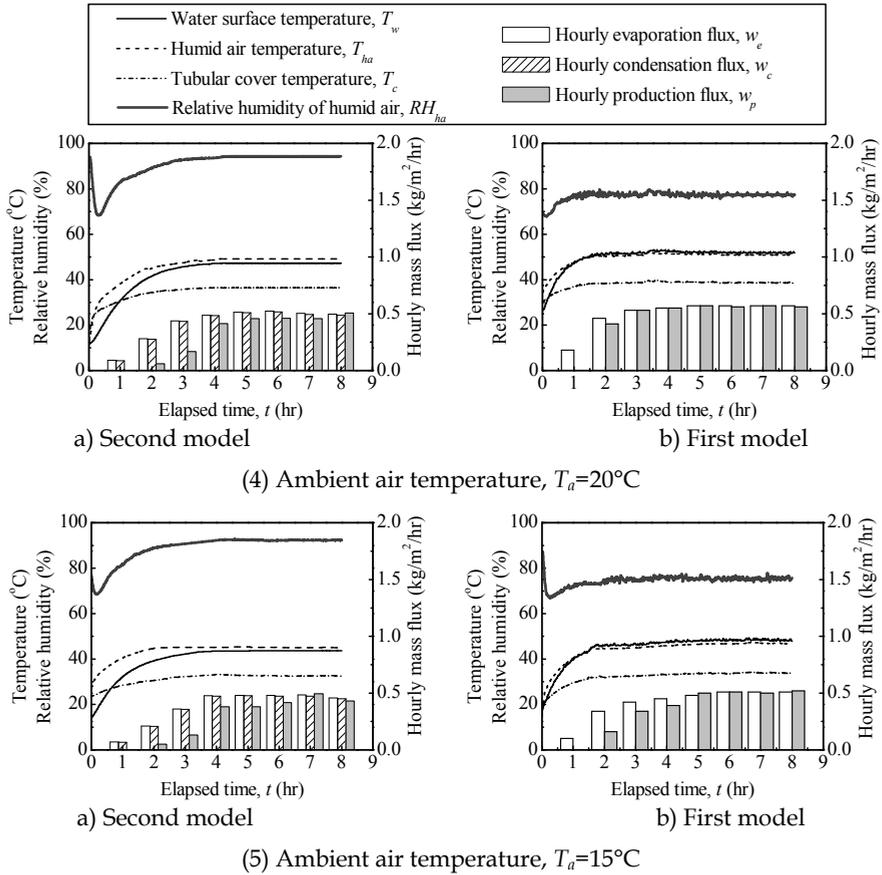


Fig. 2. Time variations of the hourly evaporation flux, w_e , hourly production flux, w_p , temperatures (T_w , T_{ha} and T_c) and RH_{ha} for different T_a ranged from 15 to 35°C for the first model and second one (Ahsan et al., 2010) (continuation)

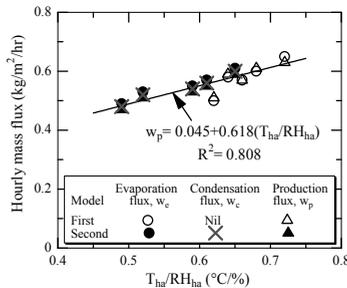


Fig. 3. Relationship between the hourly mass fluxes (w_e , w_c and w_p) and the humid air temperature and relative humidity fraction, T_{ha}/RH_{ha} , for the first model and second one (Ahsan et al., 2010)

5. Theory of mass transfer

5.1 Previous evaporation model

Islam (2006) formulated the evaporation in the TSS based on the humid air temperature and on the relative humidity in addition to the water temperature and obtained an empirical Eq. 1 of the evaporative mass transfer coefficient (m/s), h_{ew} ,

$$h_{ew} = 1.37 \times 10^{-3} + 5.15 \times 10^{-4} (T_w - T_c) \quad (1)$$

where, T_w = absolute temperature of the water surface; and T_c = absolute temperature of the tubular cover.

5.2 Purposes and research flow of present model

The main purposes and procedures of this research are as follows:

1. Making an evaporation model with theoretical expression of h_{ew}
2. Verifying the validity of the evaporation model

Three steps are taken in order to attain the two purposes described above. The purpose of the first step is to determine the value of m that is one of two unknown parameters in a new theoretical expression of h_{ew} derived by dimensional analysis. To achieve this, the evaporation experiment in this study (present laboratory-evaporation experiment) was designed and thus the correlation between the trough width, B , and hourly evaporation from the whole water surface in a trough, W , identifies the value of m (Ahsan & Fukuhara, 2008).

The purpose of the second step is to determine the value of a that is another unknown parameter in the theoretical expression of h_{ew} using the previous laboratory-TSS experimental results. Consequently, the formulization of h_{ew} is given in the second step and the first purpose is completed. Finally, the purpose of the third step is to verify the validity of the evaporation model with the new h_{ew} formulized in the second step. Therefore, the calculated evaporation mass flux was compared with the observed data obtained from the previous field-TSS experiment. Furthermore, the calculation accuracy of the previous evaporation model and another model proposed by Ueda (2000) is examined using the same field-TSS experimental data. Thus, the second purpose is achieved (Ahsan & Fukuhara, 2008).

5.3 Humid air

The density of the humid air (After Brutsaert, 1991) inside a TSS can be expressed as

$$\rho = \frac{P_o}{R_d T_{ha}} \left(1 - \frac{0.378 e_{vha}}{P_o} \right) \quad (2)$$

where, P_o = total pressure of the humid air; e_{vha} = partial pressure of water vapor in the humid air; T_{ha} = absolute temperature of the humid air; and R_d = specific gas constant of dry air. Note that $\rho = \rho_d + \rho_{vha}$, where, ρ_d = density of dry air; and ρ_{vha} = density of water vapor in the humid air. The density of the humid air on the water surface, ρ_s , can be written as (Ahsan & Fukuhara, 2008)

$$\rho_s = \frac{P_o}{R_d T_w} \left(1 - \frac{0.378 e_{vw}}{P_o} \right) \quad (3)$$

where, e_{vw} = saturated water vapor pressure. Similarly, $\rho_s = \rho_a + \rho_{vw}$, where, ρ_{vw} = density of saturated water vapor on the water surface. From Eqs. 2 and 3, the ratio of ρ to ρ_s is given by (Ahsan & Fukuhara, 2008)

$$\frac{\rho}{\rho_s} = \frac{P_o - 0.378e_{vha}}{P_o - 0.378e_{vw}} \cdot \frac{T_w}{T_{ha}} \quad (4)$$

Since the following conditions, $e_{vw} > e_{vha}$ and $T_{ha} \approx T_w$ are usually observed in a TSS (see Table 5), ρ is greater than ρ_s . This implies that the buoyancy of air occurs on the water surface and might increase the evaporation from the water surface (Ahsan & Fukuhara, 2008).

5.4 Evaporation by natural convection

We modified a diffusion equation proposed by Ueda (2000) that is applied for the evaporation from the water surface in the stagnant air with a uniform temperature. The modification of Ueda's model (present model) is attributed to the difference in the applicable condition of the diffusion equation as shown in Table 3 (Ahsan & Fukuhara, 2008).

	Present model	Ueda's model
Evaporation equation (diffusion type)	$w_x = K_m \frac{e_{vw} - e_{vha}}{\delta}$	$w_x = K_o \frac{e_{vw} - e_{vha}}{\delta}$
Physical meaning of the coefficient	K_m = Dispersion due to instability of humid air	K_o = Diffusion due to molecular motion
Air conditions on the water surface		
Temperature (°C)	Non-uniform [Upper part: low temperature, Lower part: high temperature]	Uniform
Stability of air	Unstable	Neutral

Table 3. Differences between present and Ueda's model (Ahsan & Fukuhara, 2008)

A modified diffusion equation to calculate the local evaporation mass flux, w_x , from the water surface in a trough inside a TSS is expressed as (Ahsan & Fukuhara, 2008)

$$w_x = K_m \frac{e_{vw} - e_{vha}}{\delta} \quad (5)$$

where, K_m = dispersion coefficient of the water vapor; x = transverse distance from the edge of the trough; and δ = effective boundary layer thickness of vapor pressure, e_v and depends on the convection due to the movement of the humid air in a TSS. K_m is expressed as the product of a new parameter, a_v , (Ahsan & Fukuhara, 2008) and the diffusion coefficient of water vapor in air, K_o (kg/m s Pa), i.e.

$$K_m = a_v K_o \quad (6)$$

a_v is referred to as "evaporativity" in this paper and is influenced by not only the strength of buoyancy mentioned above but also the instability of the humid air on the water surface, because the bottom boundary temperature of the humid air, T_w , is higher than the upper boundary temperature, T_c . This is the main reason why we used K_m instead of K_o , which is expressed by the following equation (Ahsan & Fukuhara, 2008),

$$K_o = \frac{DM_v}{RT_{ha}} \quad (7)$$

where, M_v = molecular weight of the water vapor; R = universal gas constant; and D = molecular diffusion coefficient of water vapor (m^2/s) at a normal atmospheric pressure and is calculated by means of the following empirical equation (After Ueda, 2000),

$$D = 0.241 \times 10^{-4} \left(\frac{T_{ha}}{288} \right)^{1.75} \quad (8)$$

Although K_o is a function of T_{ha} , the change of K_o in the range of ordinary T_{ha} is small. For example, $K_o = 1.93 \times 10^{-10} \text{ kg/m} \cdot \text{s} \cdot \text{Pa}$ for $T_{ha} = 40^\circ\text{C}$ and $2.07 \times 10^{-10} \text{ kg/m} \cdot \text{s} \cdot \text{Pa}$ for $T_{ha} = 70^\circ\text{C}$ (Ahsan & Fukuhara, 2008).

5.5 Dimensional analysis

Evaporative mass transfer is generalized by empirical equations using a dimensional analysis and correlating experimental results. Assuming that the evaporation in a TSS is induced by natural convection, the relation between δ and x is characterized using a local Grashof number, Gr , and the Schmidt number, Sc (Ueda, 2000; Ahsan & Fukuhara, 2008).

$$\frac{x}{\delta} = \frac{w_x x}{\alpha_v K_o (e_{vw} - e_{vha})} = f(Gr \cdot Sc) = a(Gr \cdot Sc)^n \quad (9)$$

The coefficient a and the power n are different for convection regimes of the humid air. The values of a and n are varied as follows (Ahsan & Fukuhara, 2008):

$a = 0.46$ and $n = 1/4$ for the laminar natural convection ($1 < Gr_B \cdot Sc < 4 \times 10^4$); and $a = 0.21$ and $n = 1/3$ for the turbulent natural convection ($4 \times 10^4 < Gr_B \cdot Sc$).

The local Grashof number is formed as a function of x :

$$Gr = \frac{gx^3}{\nu^2} \cdot \left| \frac{\rho - \rho_s}{\rho_s} \right| \quad (10)$$

where, g = gravitational acceleration; and ν = kinematic viscosity. The Schmidt number is denoted as

$$Sc = \frac{\nu}{D} \quad (11)$$

The product of the Grashof number and the Schmidt number is expressed in the form

$$Gr \cdot Sc = A \cdot \frac{gx^3}{\nu D} \quad (12)$$

where, $A = \left| \frac{\rho}{\rho_s} - 1 \right|$. Substituting Eq. 12 into Eq. 9, w_x is given by (Ahsan & Fukuhara, 2008)

$$w_x = a \alpha_v K_o (e_{vw} - e_{vha}) \left[\frac{Ag}{\nu D} \right]^n x^{3n-1} \quad (13)$$

The total evaporation mass per hour (kg/hr), i.e. hourly evaporation, W , can be obtained by integrating the local evaporation flux over the entire water surface, that is (Ahsan & Fukuhara, 2008),

$$W = 3600 \times L \times 2 \int_0^{B/2} w_x dx \quad (14)$$

where, B = width; and L = length of the trough. Integrating Eq. 14 yields the following form (Ahsan & Fukuhara, 2008):

$$W = C \alpha K_o L \left[\frac{Ag}{vD} \right]^n B^m (e_{vw} - e_{vha}) \quad (15)$$

where, $a (=aa_v)$ = evaporation coefficient; $m=3n$; and $C = \frac{3600 \times 2^{1-m}}{m}$.

When the water temperature, T_w , is different from the cover temperature, T_c , the coefficient A in Eq. 12 can be approximated by the following form (Ahsan & Fukuhara, 2008):

$$A = \frac{\rho - \rho_s}{\rho_s} \approx \beta(T_w - T_c) = \beta \Delta T \quad (16)$$

where, β = volumetric thermal expansion coefficient. Substituting Eq. 16 into Eq. 15, W is given by (Ahsan & Fukuhara, 2008)

$$W = C \alpha K_o L \left[\frac{g\beta\Delta T}{vD} \right]^n B^m (e_{vw} - e_{vha}) \quad (17)$$

Eq. 17 can be expressed in terms of the vapor density difference using the equation of state (Ahsan & Fukuhara, 2008),

$$W = C \alpha K_o L \left[\frac{g\beta\Delta T}{vD} \right]^n B^m R_v (T_w \rho_{vw} - T_{ha} \rho_{vha}) \quad (18)$$

where, R_v = specific gas constant of the water vapor. Taking into account of the fact, $T_{ha} \approx T_w$, Eq. 18 is approximated as follows (Ahsan & Fukuhara, 2008):

$$W = C \alpha K_o L \left[\frac{g\beta\Delta T}{vD} \right]^n B^m R_v \bar{T} (\rho_{vw} - \rho_{vha}) \quad (19)$$

where, $\bar{T} = \frac{T_w + T_{ha}}{2}$. Eq. 19 is transformed as (Ahsan & Fukuhara, 2008)

$$W = C \alpha K_o L \left[\frac{g\beta}{vD} \right]^n B^m R_v T^* (\rho_{vw} - \rho_{vha}) \quad (20)$$

where, $T^* = \bar{T} \Delta T^n$.

Finally, the evaporation mass flux ($\text{kg}/\text{m}^2/\text{s}$), $w(=W/3600BL)$, is calculated by the following equation (Ahsan & Fukuhara, 2008),

$$w = h_{ew}(\rho_{vw} - \rho_{vha}) \quad (21)$$

where, h_{ew} is given by (Ahsan & Fukuhara, 2008)

$$h_{ew} = \frac{2^{1-m} \alpha K_o}{m} \left[\frac{g\beta}{vD} \right]^n B^{m-1} R_v T^* \quad (22)$$

5.6 Application of the present model to the present experiment

When the vapor pressure difference, $e_{vw}-e_{vha}$, a and L are constant, Eq. 15 can be rewritten in terms of B (Ahsan & Fukuhara, 2008),

$$W = \eta B^m \quad (23)$$

where, η ($\text{kg}/\text{m}/\text{hr}$) is expressed as (Ahsan & Fukuhara, 2008)

$$\eta = C\alpha K_o L \left[\frac{Ag}{vD} \right]^n (e_{vw} - e_{vha}) \quad (24)$$

Note that e_{vha} in Eq. 24 is the vapor pressure of the stagnant ambient air surrounding the trough for the present evaporation experiment (Ahsan & Fukuhara, 2008).

Case No.	Trough length L (m)	Trough width B (m)	Radiant heat flux R_s (W/m^2)	Room air conditions	
				T_a ($^{\circ}\text{C}$)	RH_a (%)
1	0.49	0.05	Nil	29	21
2		0.10			
3		0.20			
4		0.30			
5	1.5	0.05	Nil	29	21
6		0.10			
7		0.20			
8		0.30			

Table 4. Present laboratory-evaporation experimental conditions and observed steady state values (Ahsan & Fukuhara, 2008)

6. Experiment 2: method and conditions

6.1 Present evaporation experiment

The present evaporation experiment was carried out in a temperature and relative humidity controlled room to keep e_{vw} and e_{vha} constant and the same. Table 4 shows the representative factors of the experiment such as L , B , radiant heat flux, R_s , ambient temperature, T_a , and ambient relative humidity, RH_a . The purpose is to investigate the relationship between W and B and to identify the value of m in Eq. 23. For this reason, we prepared eight troughs with four different widths (0.05, 0.1, 0.2 and 0.3m) and two different lengths (0.49 and 1.5m).

The trough was made of a corrugated carton paper of 3.0mm in thickness and covered by a black polythene film of 0.05mm in thickness. To measure the value of W , we prepared four electric balances with a minimum reading of 0.01g and each trough was placed on each electric balance. All of the electric balances were connected to computers. In this way, W was automatically and simultaneously recorded in computers at five-minute intervals. T_w was measured with a thermocouple and was recorded in a data logger. T_a and RH_a were monitored by a thermo-hygrometer (Ahsan & Fukuhara, 2008).

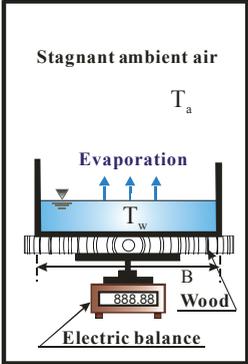
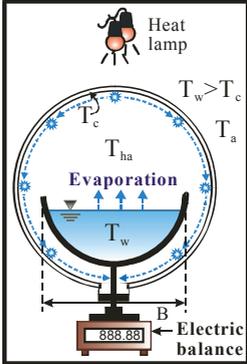
Experiment	Present evaporation experiment	Previous TSS experiment
Schematic view		
State of evaporation	From trough in stagnant air	From trough in TSS
Main differences in experimental conditions		
Ambient temperature, T_a	Constant (29°C)	Variable (20~35°C)
Ambient relative humidity, RH_a	Constant (21%)	Constant (35%)
Radiant heat flux, R_s	Nil	Variable (500~1200W/m ²)
Width of trough, B	Variable (0.05~0.3m)	Constant (0.1m)
Length of trough, L	Variable (0.49~1.5m)	Constant (0.49m)

Table 5. Laboratory experimental conditions of present and previous experiments (Ahsan & Fukuhara, 2008; Islam, 2006)

6.2 Previous TSS experiment

The results of the previous laboratory experiment using a TSS are cited to find the properties of a in Eq. 20. The TSS was comprised of a tubular cover and a black trough in it. The length and outer diameter of the tubular cover were 0.52m and 0.13m, respectively. Evaporation was enhanced with 12 infrared lamps (125W) and was controlled by changing the radiant heat (i.e. changing the height of lamp from the TSS) and the ambient air temperature (Islam, 2006).

6.3 Differences between the present and previous experiments

Table 5 summarizes the main differences between the present and previous experiments. The schematic views of both experiments are also drawn in Table 5. The size of the trough was changed in the present evaporation experiment, but the external environment

surrounding the trough maintained the same conditions. On the other hand, the external conditions (R_s and T_a) were changed in the previous experiment, but the same trough size ($L=0.49\text{m}$ and $B=0.1\text{m}$) was used then (Ahsan & Fukuhara, 2008).

6.4 Previous field-TSS experiment

In order to support the validity of the present model, the previous field experimental results are cited in this paper. The same specification of TSS was produced for both laboratory and filed experiments (Islam, 2006).

7. Relation between W and B

T_w for the eight different troughs were nearly the same (maximum difference 0.5°C) and T_a and RHa were also the same (29°C and 21% , respectively). Therefore, it was established that $evw=evha$ was the same for every experimental case. Furthermore, it is assumed that the instability of air and the strength of buoyancy on the water surface might be the same for every experimental case. Therefore, we expected that α would have the same value for every experimental case and that η is treated as a constant (Ahsan & Fukuhara, 2008).

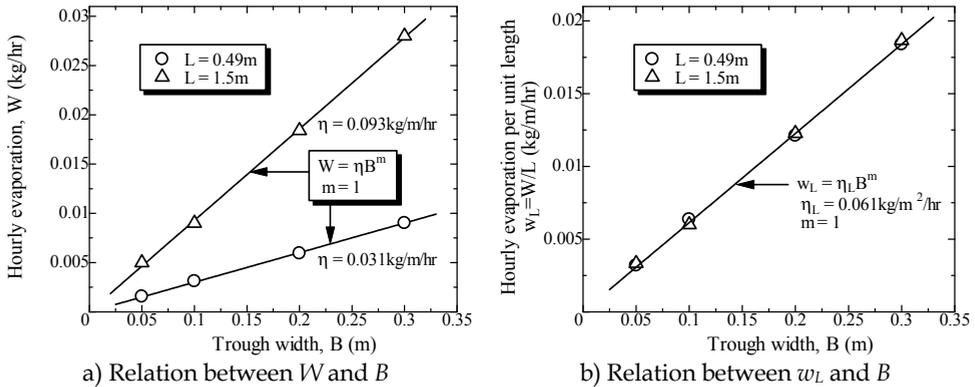


Fig. 4. Variation of the hourly evaporation by changing the trough width and length (Ahsan & Fukuhara, 2008)

The values of $Gr_B \cdot Sc \left(= A \cdot \frac{gB^3}{\nu D} \right)$ ranged from 4.03×10^4 to 7.70×10^6 . The state of air flow over

the trough would be, therefore, in turbulent natural convection (Ahsan & Fukuhara, 2008).

Fig. 4(a) shows the effect of the trough size (B and L) on W . W is linearly proportional to B . We found that the value of m in Eq. 23 is 1, i.e. $n=1/3$, regardless of L . W for $L=1.5\text{m}$ is nearly three times larger than that for $L=0.49\text{m}$ for the same B . Consequently, the hourly evaporation per unit length, $w_L (=W/L)$, is expressed as a function of B as shown in Fig. 4(b) and all data is on a regression straight line; $w_L = \eta_L B^m$ where $\eta_L (= \eta/L)$ is $0.061\text{kg/m}^2/\text{hr}$ and might be independent of L for $0.49 \leq L \leq 1.5\text{m}$ (Ahsan & Fukuhara, 2008). The value of m is 1 and is in agreement with the results of Ueda (2000).

Using $\eta_L = 0.061\text{kg/m}^2/\text{hr}$ and $m=1$, a can be calculated by Eq. 24. It can be observed that the value of a is a constant ($=0.06$) for every experimental case, regardless of B (Ahsan & Fukuhara, 2008).

8. Evaporation coefficient

The results of the previous laboratory-TSS experiment under twelve sets of external conditions are quoted here (Islam, 2006). Since the vapor density difference, $\rho_{vw} - \rho_{vha}$, is different for every experimental case unlike the present evaporation experiment, a should be calculated by Eq. 25 after substituting $m=1$ into Eq. 20 (Ahsan & Fukuhara, 2008),

$$\alpha = \frac{W}{3600K_oL \left[\frac{g\beta}{\nu D} \right]^{\frac{1}{3}} BR_v T^* (\rho_{vw} - \rho_{vha})} \quad (25)$$

As $GrB : Sc$ exceeds 4×10^4 for every case, it is inferred that the humid air flow on the trough in the TSS would be in turbulent natural convection state (Ahsan & Fukuhara, 2008).

The temperature difference, $T_w - T_c$, might be one of the parameters that represent the instability of the humid air. Since T_w is higher than T_c , it is inferred that the humid air would become unstable as the temperature difference $T_w - T_c (>0)$ increases. Based on this concept, Fig. 5 shows the relation between $T_w - T_c$ and a . The value of a is proportional to $T_w - T_c$ and the regression can be expressed as (Ahsan & Fukuhara, 2008)

$$\alpha = 0.123 + 0.012(T_w - T_c) \quad (26)$$

Substituting Eq. 26 and $m=1$ into Eq. 22, h_{ew} is given by (Ahsan & Fukuhara, 2008)

$$h_{ew} = \left[0.123 + 0.012(T_w - T_c) \right] \cdot \left[\frac{g\beta}{\nu D} \right]^{\frac{1}{3}} K_o R_v T^* \quad (27)$$

The hourly evaporation mass flux, $w_h (=W/BL)$, is expressed as (Ahsan & Fukuhara, 2008)

$$w_h = 3600h_{ew}(\rho_{vw} - \rho_{vha}) \quad (28)$$

Once the four parameters (T_w , T_{ha} , T_c and RH_{ha}) are measured, w_h can be calculated by combining Eqs. 27 and 28 (Ahsan & Fukuhara, 2008).

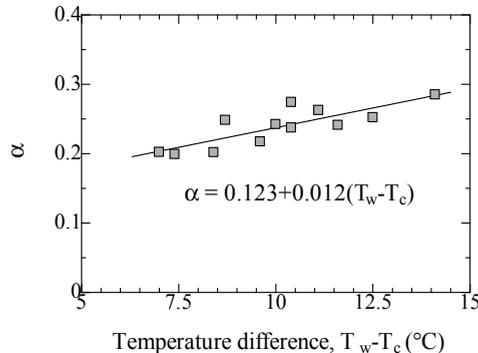


Fig. 5. Relation between the evaporation coefficient, a , and the temperature difference, $T_w - T_c$, obtained from the previous laboratory-TSS experiment (Ahsan & Fukuhara, 2008)

9. Model validation

The applicability of the present and previous evaporation models were examined by comparing them with the previous field-TSS experimental results (Islam, 2006) obtained in Fukui, Japan (September 29 and October 6, 2005).

Fig. 6(a) and (b) show the calculation accuracy of w_h calculated by the two models (present and previous) and Ueda's model. The accuracy of the present model is satisfactory and is applicable to both laboratory and field experiments. However, w_h calculated by the previous model using the empirical Eq. 1 slightly underestimates the observed w_h (Ahsan & Fukuhara, 2008).

Ueda's model also underestimates the calculated value and the deviation from the observed value is largest among the three models. Using the coefficient K_o related to the molecular diffusion might be the reason for such underestimation. A better estimation of w_h could be found (Ahsan & Fukuhara, 2008) using Ueda's model when $a_v (=K_m/K_o)$ is 1.14 (in average), assuming that the coefficient a in Eq. 9 is 0.21 for turbulent natural convection according to Ueda (2000).

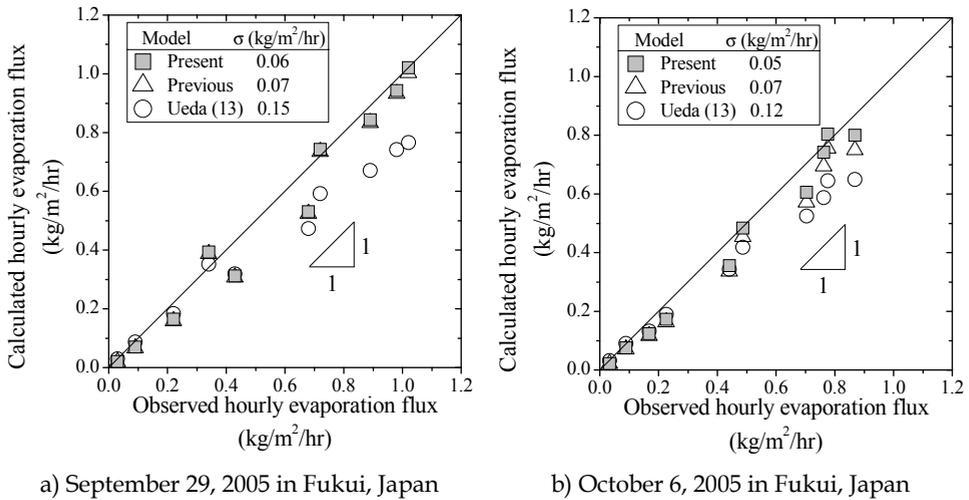


Fig. 6. Comparison of calculated hourly evaporation mass flux with the observed value of the previous field-TSS experiment (Ahsan & Fukuhara, 2008)

The calculation accuracy of these three models was quantitatively evaluated by the root mean squared deviation, σ . That is (Ahsan & Fukuhara, 2008),

$$\sigma = \sqrt{\left[\frac{1}{N} \sum_{i=1}^N (w_{hoi} - w_{hci})^2 \right]} \quad (29)$$

where, w_{hoi} = observed hourly evaporation mass flux; w_{hci} = calculated hourly evaporation mass flux; and N is the data number. The value of σ is given for each model in Fig. 6. The present model has the smallest σ among the three models and the difference in σ between two models (present and previous) is small. σ of Ueda's model is more than twice than that of the present model (Ahsan & Fukuhara, 2008).

10. Conclusion

The cover material of the first model of the Tubular Solar Still (TSS), a transparent vinyl chloride sheet was changed to a polythene film for the second model. Thus, the second model is simpler, lighter, cheaper and more durable than the first one. These improvements make the assembly and maintenance of the new TSS easier. A special experimental technique was developed to observe the evaporation, condensation and production performance independently and simultaneously. As a result, the evaporation was detected first and then the condensation and the production followed it in turn. As for second model, the hourly evaporation and production fluxes were slightly lower than the first one under the same experimental conditions. It was revealed that the relative humidity of the humid air was definitely not saturated and the hourly evaporation, condensation and production fluxes were proportional to the humid air temperature and relative humidity fraction (Ahsan et al., 2010).

An evaporative mass transfer model was presented with a semi-theoretical expression of the evaporative mass transfer coefficient for a TSS using the dimensional analysis taking account of the humid air properties inside the still. Findings revealed from the present laboratory-evaporation experimental results that the hourly evaporation is linearly proportional to the trough width, B , regardless of the trough length, L , for $0.49 \leq L \leq 1.5\text{m}$ (Ahsan & Fukuhara, 2008). The movement of the humid air in the TSS belongs to turbulent natural convection state. The evaporation coefficient is proportional to the temperature difference between the water in a trough and the tubular still cover. The present model was able to reproduce the hourly evaporation mass flux obtained from the previous field-TSS experiment. It is concluded that once the four parameters (Ahsan & Fukuhara, 2008); that is, the water temperature, humid air temperature, tubular cover temperature and the relative humidity of humid air are measured, the present model is capable of evaluating the diurnal variation of evaporation mass flux from the water surface in a trough with an arbitrary size.

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12. Nomenclature

B	=	width of the trough (m);
D	=	molecular diffusion coefficient of water vapor (m^2/s);
e_v	=	vapor pressure (Pa);
e_{vha}	=	partial pressure of water vapor in the humid air (Pa);
e_{vw}	=	saturated water vapor pressure (Pa);
g	=	gravitational acceleration (9.807 m/s^2);
Gr	=	Grashof number (-);

h_{ew}	=	evaporative mass transfer coefficient from water surface to humid air (m/s);
K_m	=	dispersion coefficient of the water vapor (kg/m s Pa);
K_o	=	diffusion coefficient of the water vapor (kg/m s Pa);
L	=	length of the trough (m);
M_o	=	molecular weight of the water vapor (18.016 kg/kmol);
P_o	=	total pressure in the humid air (101325 Pa);
R	=	universal gas constant (8315 J/kmol K);
R_d	=	specific gas constant of dry air (287.04 J/kg K);
R_s	=	radiant heat flux (W/m ²);
R_v	=	specific gas constant of the water vapor (461.5 J/kg K);
RH_a	=	relative humidity of ambient air (%);
RH_{ha}	=	relative humidity of humid air (%);
Sc	=	Schmidt number (-);
T_a	=	ambient air temperature (K);
T_c	=	tubular cover temperature (K);
T_{ha}	=	humid air temperature (K);
T_w	=	water surface temperature (K);
\bar{T}	=	mean temperature of water and humid air (K);
w	=	evaporation mass flux (kg/m ² s);
w_h	=	hourly evaporation mass flux (kg/m ² hr);
w_{hci}	=	calculated hourly evaporation mass flux (kg/m ² hr);
w_{hoi}	=	observed hourly evaporation mass flux (kg/m ² hr);
w_L	=	hourly evaporation per unit length (kg/m hr);
w_x	=	local evaporation mass flux (kg/m ² s);
W	=	hourly evaporation (kg/hr);
x	=	transverse distance from the edge of trough (m);
a	=	evaporation coefficient (-);
a_v	=	evaporativity (-);
β	=	volumetric thermal expansion coefficient (1/K);
δ	=	effective boundary layer thickness of vapor pressure (m);
ΔT	=	temperature difference between water surface and cover (K);
ν	=	kinematic viscosity (m ² /s);
ρ	=	density of humid air (kg/m ³);
ρ_d	=	density of dry air (kg/m ³);
ρ_s	=	density of humid air on the water surface (kg/m ³);
ρ_{vha}	=	density of water vapor in the humid air at T_{ha} (kg/m ³);
ρ_{vw}	=	density of saturated water vapor on the water surface at T_w (kg/m ³); and
σ	=	root mean squared deviation (kg/m ² hr).

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